

A deep reservoir of primordial helium in the Earth

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The Earth has a layered internal structure with the crust, upper mantle, mantle transition zone, lower mantle, outer core, and inner core from the surface to the center. In the Earth's formation stage at approximately 4.6 billion years ago, the heavy metal components were separated from silicates and sank in the magma ocean, and a core formed at the center of the Earth. In this core-mantle separation process, partitioning of noble gases between the core and mantle



occurred. Credit: Taku Tsuchiya, Ehime University

Noble gases, including helium, neon, and argon, are characterized by high chemical inertness which causes low reactivity with other materials and high volatility. Among them, ³He, ²⁰Ne, and ³⁶Ar are particular isotopes which were components of the primordial solar nebula existing in space before the Earth had formed. ³He is also known to have been produced by the Big Bang and a substantial amount is contained in ocean island basalts, e.g., Loihi Seamount, Hawaii (e.g., Dixon et al., 2000). Such basalts are hot spot rocks which originated in the Earth's deep interior, indicating that ³He was stored somewhere in the deep Earth. It is surprising that such primordial helium has been confined in the Earth's interior for 4.6 billion years, from the time of the Earth's formation to now, even though noble gases are highly volatile. Considering the vigorous mantle convection throughout the geological time scale (e.g., Van der Hilst et al., 1997; Wang et al., 2015), it would seem unlikely that noble gases would be trapped inside the Earth so long. Although it has been suggested that the candidates for the location of the reservoir of primordial helium are the deepest mantle and the core (image 1), its location remains unclear. This is one of the biggest mysteries in deep Earth science and still under intense debate.

The <u>outer core</u>, composed mainly of <u>liquid iron</u>, is a candidate for the reservoir of primordial <u>helium</u>, and there is a possibility that helium is supplied from this area to the mantle. Such <u>noble gases</u> could be carried up to the surface with upwelling mantle plumes. This seems a reasonable scenario to explain the fact that rocks collected in the active hot spot areas, such as in Loihi Seamount and Iceland, contain high concentrations of primordial noble gases. If the outer core is the reservoir of noble gases, the necessary amounts would have to be dissolved in liquid iron under <u>high pressure</u>. However, previous



experimental studies reported that at relatively low pressures from 1 atm to 20 GPa, noble gases generally prefer silicates (the mantle) to metals (the core) (e.g., Bouhifd et al., 2013). (The property by which a particular solute is dissolved into different coexisting solvents in different amounts is called element partitioning.) On the other hand, there exists no study so far which has investigated the property of metal/silicate partitioning of noble gases at pressures of 10 GPa to 100 GPa, corresponding to the conditions where the Earth's proto core reacted with the magma ocean in the early stage of the Earth's formation. Therefore, it is hard to exclude the possibility that the core is a reservoir of noble gases. If noble gases change to prefer metals with increasing pressure (a property called siderophile), more could be dissolved into the core, and it is important to clarify the partitioning properties of noble gases.

Precise experimental measurements of elements partitioning under high pressure are quite difficult, so in this study, by means of the quantum mechanical computer simulation technology called the ab initio method, the partitioning properties of helium and argon between liquid iron and molten silicate (magma) were investigated in the wide pressure range of 20 GPa to 135 GPa. Computer simulations of element partitioning were conducted by calculating the reaction energies when noble gases are dissolved into liquid iron and molten silicate. By comparing these reaction energies, the relative differences in the equilibrium of the noble gas concentrations in coexisting liquid iron and molten silicate could be estimated. Based on the fundamental principle of thermodynamics, noble gases are dissolved more into a solvent with a smaller reaction energy, and thus larger differences in the reaction energies more greatly enhance the contrast in the noble gas concentrations in liquid iron and molten silicate. Special techniques are required to compute the reaction energies of noble gases with liquids such as liquid iron and molten silicate. In this study, this was conducted by combining a method called the thermodynamic integration method, authorized by statistical



mechanics, with the ab initio molecular dynamics method.



The thermodynamic indication showing how much solute element (helium in this study) is dissolved in a coexisting solvent (liquid iron and molten silicate in this study) is named the partitioning coefficient. The partitioning coefficient, represented by a ratio of concentrations of the solute in the respective solvent, becomes 100 (=1) when solubility of the solute is equal in both solvents. In this study, helium is preferentially partitioned into liquid iron (molten silicate) when it is greater (smaller) than 100. The metal/silicate partitioning coefficient of helium has been measured experimentally in the relatively lower pressure range only, but it was determined in the entire mantle pressure range in this study. Partitioning coefficients determined at 20 GPa, 60 GPa, and 135 GPa are not very different and around 10-2, indicating that pressure has a low effect on the metal/silicate partitioning of helium. This value means that helium is dissolved in



molten silicate at around 100 times more than in liquid iron. Credit: Taku Tsuchiya, Ehime University

The calculations of the partitioning properties of noble gases between liquid iron and molten silicate obtained by these original techniques indicate for the first time in the world that noble gases remain, preferring molten silicate to liquid iron up to the core-mantle boundary pressure (135 GPa), and there is no distinct increase in their siderophility. The amount of helium dissolved in the core in the early stage of the Earth's formation is considered to be approximately 1/100 of the amount dissolved in the mantle (image 2). (In contrast, argon is found to become more siderophile with increasing pressure. The different high-pressure behaviors are caused by the different atomic sizes of helium and argon.) This result, showing no considerable pressure effects, suggests that the core is unsuitable as the primordial reservoir, but the estimated total amount of ³He stored in the present hot spots.

Even though 100 times more helium was dissolved into the magma ocean, most of it would have evaporated into the air while it solidified and only marginal amounts would be left due to its high volatility. In contrast, helium dissolved in the core during the proto core formation in the magma ocean was confined to the core after the magma ocean solidified. It is considered that such helium has been gradually seeping into the mantle across the core-mantle boundary and ascending to the surface with upwelling plumes over a long period of time. It can be measured in the hot spot rocks even now. These results provide conclusive support showing that the ³He reservoir is at the <u>core</u>. This is an important insight for the location of the primordial reservoir, one of the long-standing mysteries in geoscience.



More information: Zhihua Xiong et al. Helium and Argon Partitioning Between Liquid Iron and Silicate Melt at High Pressure, *Geophysical Research Letters* (2020). DOI: 10.1029/2020GL090769

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